## NH<sub>3</sub>, H<sub>2</sub>S, and the Radio Brightness Temperature Spectra of the Giant Planets

Thomas R. Spilker

Jet Propulsion Laboratory, California Institute of Technology

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Recent radio interferometer observations of Neptune enable comparisons of the radio brightness temperature (T<sub>B</sub>) spectra of all four giant planets. Figures 1, 2, 3, and 4 show the T<sub>B</sub> spectra of Jupiter, Saturn, Uranus, and Neptune, respectively, from 0.1 10 tens of cm wavelength. The data shown are collected from many observers. Data for Jupiter, Saturn, and Uranus are those cataloged by de Pater and Massie [1985], plus the Saturn VI. A data by Grossman *et al.* [1989]. Figure 3, Uranus, shows only data acquired since 1973. Before 1973. Uranus' T<sub>B</sub> increased steadily as its pole moved into view, causing significant scatter in those data. Neptune data at>1cm, all taken at the VI. A, are collected from de Pater and Richmond [1989], de Pater *et al.* [1991], and Hofstadter [1993]. For a variety of reasons, siagle-dish data at those wavelengths are muchnoisier than the more reliable VI. A data and have been ignored. Single-dish data by Griffin and Orton [1993] shortward of 0.4 cm are shown, along with the OVRO datum at 0.266 cm by Muhleman and Berge [1991].

Spectra of Jupiter, Saturn, and Neptune share certain gross characteristics. In each case T<sub>B</sub> at 0.1 cm is within -30 K of that at 1.3 cm, which is in the 120-140 K range. The spectra increase monotonically with wavelength only longward of 1.3 cm. Ammonia (N1 I<sub>3</sub>), whose strong inversion spectrum peaks at -1.3 cm, is known to be an important tropospheric constituent at Jupiter and Saturn. Its signature on the Jovian spectrum is obvious, causing the prominent "hole" at 1.3 cm. At Saturn it is a bit more subdued but is the source of that spectrum's change in slope at 1,3 cm. Radiative transfer models of Jupiter and Saturn using near-solar deep NH<sub>3</sub> abundances agree well with the data [de Pater, 1990].

Uranus' 'T<sub>a</sub> spectrum dots not fit this pattern. It increases monotonically with wavelength over the entire range shown in figure 3, with no evidence of a break in slope near 1.3 cm. T<sub>B</sub> is -175 K at 1.3 cm, -80 K warmer than at 0.1 cm and much warmer than the other three planets. At -20 cm and 0.1-0.4 cm Uranus' T<sub>B</sub> are quite close to Neptune's, but in the 1-10 cm range Uranus averages 30-50 K colder than Neptune. Gulkis *et al.* [1978] first showed that Uranus radiative transfer models with near-solar NI 1<sub>3</sub> deep abundances predict '1'<sub>a</sub> at cm wavelengths that are much too cold. Using an NI 1<sub>3</sub> abundance about 1% of solar fit the data best, but far from pert'ectly. They offered one possible cause for the apparent NH<sub>3</sub> depletion: a superabundance of 112S could react out most of the NI 1<sub>3</sub>. Note the H<sub>2</sub>S was postulated. "1'here is no direct observational evidence of 11<sub>2</sub>S, which has not yet been detected at any of the giant planets. It was merely considered the most likely candidate to deplete NI 1<sub>3</sub>.

Some researchers suggest a significant 1 I<sub>2</sub>S superabundance at Neptune also. Problems fitting radiative transfer models to cm data prompted dePater *et al.* [ 1991] to invoke NII<sub>3</sub>-depleting H<sub>2</sub>S at Neptune, and to suggest that H<sub>2</sub>S

might contribute significantly 10 the total opacity. Recently DeBoer and Steffes [1994] (hereafter DBS) made lab measurements of cm 1  $I_2$ S opacities and found them a factor of two larger than Van Vleck-Weisskopf predictions. Based on this they suggest  $11_2$ S may be the major source of cm opacity in Neptune's upper troposphere, and reinterpret Lindal's [1992] Voyager 2 radio occultation data. Lindal assumed all opacity at the 6.3 bar level, the deepest probed, was due 10 NH<sub>3</sub> and derived a number mixing ratio of  $5 \times 10^{-7}$ . DBS assume all opacity there is due to  $H_2$ S and derive a mixing ratio of  $1.7 \times 10^4$ . For support they compare radiative transfer model results m the  $T_B$  data, but rely heavily on the very noisy single-dish data at > 1 tin, where the VLA data are much more reliable.

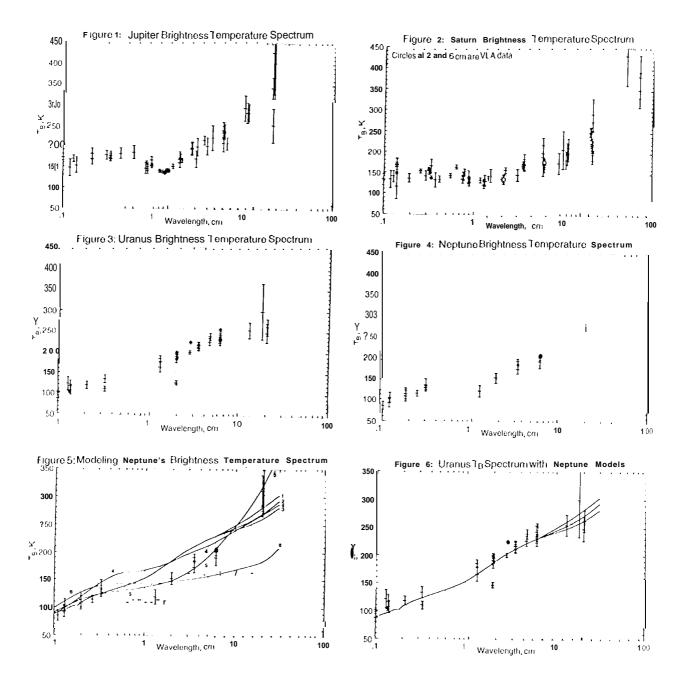
Figure 5 duplicates Figure 4 except it includes results from various radiative transfer models. Models 1-4 arc after DBS, with 30 times solar 1120 and CH<sub>4</sub>; NH<sub>3</sub> and 11<sub>2</sub>S abundances, respectively, are 0.5 solar and 15 times solar for model 1, solar and 18 times solar for model 2, twice solar and 25 times solar for model 3, and solar and 6 times solar for model 4. Models 1-3, which are identical shortward of 6 cm, yield 1.7 x 10 H<sub>2</sub>S above the NH<sub>4</sub>SH cloud, while model 4 yields Lindal's 5 x 10<sup>7</sup> NH<sub>3</sub>. Model 5 is after de Pater and Richmond [1989], using an ~2% solar NH<sub>3</sub> mixing ratio (3 x 10<sup>-6</sup>) throughout the atmosphere, limited by saturation. Model 6, by the author, uses approximately solar NH<sub>3</sub> (2 x 10<sup>-4</sup>) and no H<sub>2</sub>S to demonstrate that Neptune models with uniformly near-solar NH<sub>3</sub> abundances are inconsistent with the observed spectrum.

Only the models with NH<sub>3</sub> above the NI14S11 cloud reproduce Neptune's  $T_B$  dip at cm wavelengths. Model 5, with more NH<sub>3</sub> than model 4, provides the best fit; even more NH<sub>3</sub> would provide a better fit, further decreasing  $T_B$  shortward of 1 tin. This dots not conflict with Lindal's result, since he stales NH<sub>3</sub> is probably still saturated at the deepest datum. Due to the  $11_2$ S spectrum's simple  $f^{-2}$  dependency longward of 0.4 cm, models dominated by H<sub>2</sub>S above the NH<sub>4</sub>S11 cloud (DBS models 1-3) deviate <10 K from a straight line on the plot, quite unlike the data. Reproducing the  $T_B$  dip with such an absorber requires a relatively this tropospheric layer with a much larger absorber mixing ratio than adjacent layers. Since there is no viable mechanism to maintain such a layer, it is highly unlikely that the observed cm opacity in Neptune's upper troposphere is primarily due (o 1 I<sub>2</sub>S. Thus Neptune's radio spectrum requires NH<sub>3</sub>, or another species with an opacity peak near 1-2 cm; in the upper troposphere.

Applying the DBS models 1-3 to Uranus leads to a different conclusion for that planet. Upper tropospheric T-P (temperature-pressure) relations for Uranus and Neptune arc very similar; all equal pressures, their temperatures differ by ~5K at most from well above their tropospheres to the deepest level probed by radio occultation [1 indal *et al.*, 1987;Lindal, 1992]. Given a fixed set of constituent abundance profiles, a model using Neptune's T-P profile will

yield a  $T_B$  spectrum quite similar to one produced using Uranus '1'-[' profile. The dissimilarity of the two planets' observed radio spectra makes it unlikely they have similar constituent profiles. Figure 6 shows the result of using DBS'  $1 I_2 S$ -dominated models of Neptune as first approximations to such models for Uranus. The models fit Uranus' observed spectrum much better than Neptune's, suggesting that tropospheric constituents whose cmopacities have  $f^2$  dependencies, such as  $II_2 S$ , are sufficient to explain Uranus' radio  $T_B$  spectrum. Thus, while Neptune seems to need a small (relative to solar') but nontrivial amount of N113 inits upper troposphere, Uranus does not.

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